



Assessment of Climate Change Impact on Summer Low Flows in Mountainous Regions using a Coupled Hydrology-Glacier Model

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Left to Parker Ridge" ~ Randy Hayashi Paintings





Applying the Variable Infiltration Capacity model



- ✓ Large Spatial Scales
- ✓ Physical Representation of *Most* Major Processes
- X Not Glaciers





Location of Glaciers in the Study Area



~ Randolph Glacier Inventory (RGI; Pfeffer et al. 2014)





Modelling in Mountainous-Glaciated Basins



Observed Streamflow 1985-1995





VIC Limitations

- 1. No explicit treatment of glacier mass or glacier energy balance.
- 2. Naïve representation of sub-grid land cover and topographic variability.
- 3. No lateral fluxes between grid cells.





Implementing a Glacier Mass Balance Model in VICGL

The model estimates the rate of accumulation and ablation of glaciers, or mass balance, which serves two purposes:

- 1. To directly calculate ice melt contribution to runoff in VIC.
- 2. To act as a forcing to the (offline) dynamics model.





Glacier Mass and Glacier Energy Balance Model in VICGL



Predominantly based on model of Klok and Oerlemans (2002):

$$F = S_{in}(1 - \alpha) + L_{in} + L_{out} + Q_s + Q_L + Q_r = Q_m + G$$

where:

- S_{in} = incoming solar radiation L_{in} = incoming longwave radiation L_{out} = outgoing longwave radiation Q_s = sensible heat Q_L = latent heat
- Q_r = heat advected from rainfall T_r Q_m = energy for ice melt P_r G = glacier heat flux T_r L_s = latent heat of sublimation T_r L_m = latent heat of melt d_{surf} = thickness of glacier surface layer
- T_{surf} = glacier surface temperature p_i = density of ice
 - c_i = heat capacity of ice
 - α = glacier albedo

Adapted from Schnorbus et al. 2016





VIC versus VICGL Water Balance



P E P E S_{snow} P E S_{glac} S_{glac} S_{soil}

R

Revised Water Balance (VICGL)

$$R = P - E - \Delta S_{soil} - \Delta S_{snow} - \Delta S_{glacier}$$

Adapted from Schnorbus et al. 2016





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Refactoring VIC to VICGL: from Bands to HRUs

Bull River near Wardner, BC



Adapted from Schnorbus et al. 2016





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The Regional Glaciation Model – Clark et al. 2016

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Figure 1 | Study region and subregions in the Canadian Cordillera of western Canada. Present-day (2005) glacier extent is indicated in white. The yellow rectangle indicates the location of Columbia Reach drainage basin. Inset: Study region (black rectangle) within northwestern North America.



Figure 2 | Comparisons of observed and modelled ice hypsometry for reference year 2005. For each of the ten subregions, pairs of bell-shaped curves show the distribution of ice area (normalized to the largest observed area) with elevation for the observed (green) and modelled (blue) ice extent.

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Projected deglaciation of western Canada in the twenty-first century

LETTERS

Garry K. C. Clarke^{1*}, Alexander H. Jarosch², Faron S. Anslow³, Valentina Radić¹ and Brian Menounos⁴

Retreat of mountain glaciers is a significant contributor to sea-level rise and a potential threat to human populations through impacts on water availability and regional hydrology. Like most of Earth's mountain glaciers, those in western North America are experiencing rapid mass loss¹². Projections of future large-scale mass change are based on surface mass balance models that are open to criticism, because they ignore or greatly simplify glacier physics. Here we use a high-resolution regional glaciation model, developed by coupling physics-based ice dynamics with a surface mass balance model, to project the fate of glaciers in western Canada. We use twenty-first-century climate scenarios from

and global scales^{5,6,8,15}. A common feature of these models is that they lack a physics-based treatment of glacier dynamics. The central contribution of our study is thus to simulate the changes in ice thickness and extent over a large region using a high-resolution model of glacier dynamics, which yields year-to-year changes in ice area and volume for the entire study region.

Our study area is Alberta and British Columbia (BC) in western North America (Fig. 1), where glaciers account for an estimated area of 26,700 km² (ref. 18) and volume of 2,980 km³ (ref. 19). The geographical scale is comparable to that of other glacierized mountain regions, such as South America (~31,900 km²; ref. 7), the Himalaya and Karakoram (~22,800 km² and ~18.000 km²;





The Regional Glaciation Model – Clark et al. 2016



Clarke, G. K. C., Jarosch, A. H., Anslow, F. S., Radic, V., and Menounos, B.: Projected deglaciation of western Canada in the twenty-first century, Nat. Geosci., 8, 372–377, 2015.





VICGL HydroConductor RGM

VICGL



Run 3 Hourly with Daily Output

HydroConductor from VICGL to RGM

- VIC grid cell to RGM grid cell

remapping

RGM requires

- Surface DEM
- Bed DEM
- Mass balance

RGM produces

- An updated surface DEM

HydroConductor from RGM to VICGL

- Turns this updated glacier mask into updated HRUs for VICGL by modifying the:
 - State file
 - Vegetation Parameter file
 - Snow Band file

RGM



Run Monthly with Annual Output



- **1.** Explicit treatment of glacier mass or glacier energy balance.
- 2. Reasonable representation of sub-grid land cover and topographic variability.
- 3. Lateral <u>ice</u> fluxes between grid cells.







Thank you

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"Left to Parker Ridge" ~ Randy Hayashi Paintings



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